

Seismic imaging with generalized Radon transforms: stability of the Bolker Condition

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Joint work with Christine Grathwohl, Peer Kunstmann, and Andreas Rieder
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An inverse problem for the acoustic wave equation

$u(t; \mathbf{x}, \mathbf{x}_s)$ acoustic potential at $\mathbf{x} \in \mathbb{R}^d$, $d \in \{2, 3\}$, at time $t \geq 0$

$$\frac{1}{v_p^2} \partial_t^2 u - \Delta_{\mathbf{x}} u = \delta(\mathbf{x} - \mathbf{x}_s) \delta(t)$$

$v_p = v_p(\mathbf{x})$ speed of sound, \mathbf{x}_s excitation (source) point.

Seismic imaging of
backscattered (reflected) fields

$u(t; \mathbf{x}_r, \mathbf{x}_s)$, $t \in [0, T_{\max}]$,
 $(\mathbf{x}_r, \mathbf{x}_s) \in \mathcal{R} \times \mathcal{S}$ where \mathcal{R}/\mathcal{S}
sets of receiver/source points,
and T_{\max} observation period.

Simplifying ansatz

Consider the ansatz

$$\frac{1}{v_p^2(\mathbf{x})} = \frac{1 + n(\mathbf{x})}{v^2(\mathbf{x})},$$

with v_p the actual sound speed (high frequency content n) and $v = v(\mathbf{x})$ smooth and known background velocity, $v_p \sim v$.

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Relation between velocity and travel time:

$$|\nabla_{\mathbf{x}} \tau| = v^{-1} \quad \tau(\mathbf{x}_s, \mathbf{x}_s) = 0.$$

In general, you know v and need to find τ

Determining n using the Ansatz

To determine n , we use a reduction from the solution to the wave equation to an expression for n (e.g. SYMES 1998):

$$Fn(t; \mathbf{x}_r, \mathbf{x}_s) = \int \frac{n(\mathbf{x})}{v^2(\mathbf{x})} A(\mathbf{x}, \mathbf{x}_s) A(\mathbf{x}, \mathbf{x}_r) \delta(t - \tau(\mathbf{x}_s, \mathbf{x}) - \tau(\mathbf{x}, \mathbf{x}_r)) d\mathbf{x}$$

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Note: F is a Radon transform that, for each t , integrates n over reflection isochrones (surfaces of constant travel time t from \mathbf{x}_s to \mathbf{x} to \mathbf{x}_r): $t = \tau(\mathbf{x}_s, \mathbf{x}) + \tau(\mathbf{x}, \mathbf{x}_r)$.

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Under certain conditions F is an Fourier integral operator.

Microlocal and other properties of F and reconstruction methods in various geometric settings have been studied by many authors, such as:

BEYLKIN 1985, RAKESH 1988, NOLAN/SYMES 1997, ... etc.

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A series of works with Christine Grathwohl, Peer Kunstmann, and Andreas Rieder (working with geophysicists at KIT)

We now talk about some of our recent results!

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- ▶ *Common offset data acquisition geometry*: Let $\alpha \geq 0$ be the common offset. Then, sources and receivers on the surface are parameterized by $s \in \mathbb{R}$ and

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Some geophysicists approximate known background velocity in layered media by an affine speed.

The 2D situation for *affine* velocity $v(x_1, x_2) = b + ax_2$
for a and b positive

Travel time: $\tau((r, 0)(x_1, x_2)) = \frac{1}{a} \operatorname{acosh} \left(1 + \frac{a^2}{2b} \frac{(x_1-r)^2 + x_2^2}{b+ax_2} \right)$

See Lesson 41 in [Slotnick, Lessons in Seismic Computing, 1959].

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Domains: $X = \{\mathbf{x} \in \mathbb{R}^2 : x_2 > x_{\min}\}$ and $Y = S \times]t_{\min}, \infty[$ where S is an open subset of \mathbb{R} and

$$x_{\min} := \frac{b}{a} \left(\sqrt{1 + \frac{a^2 \alpha^2}{b^2}} - 1 \right), \quad t_{\min} := \frac{2}{a} \operatorname{asinh} \left(\frac{a\alpha}{b} \right),$$

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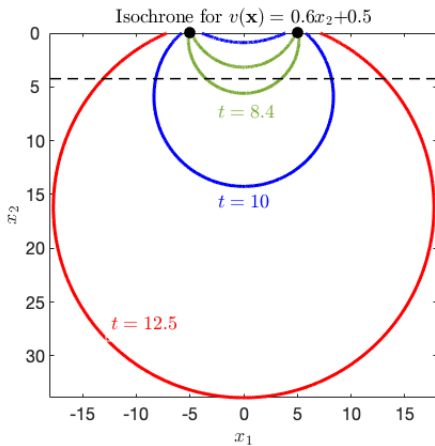
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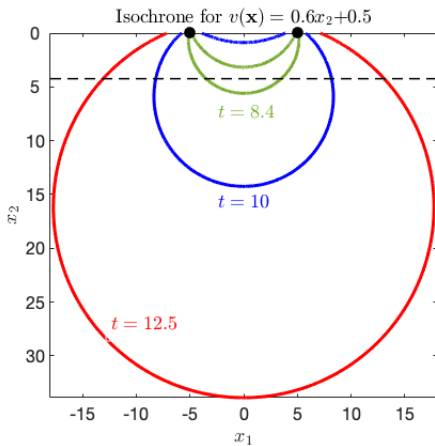
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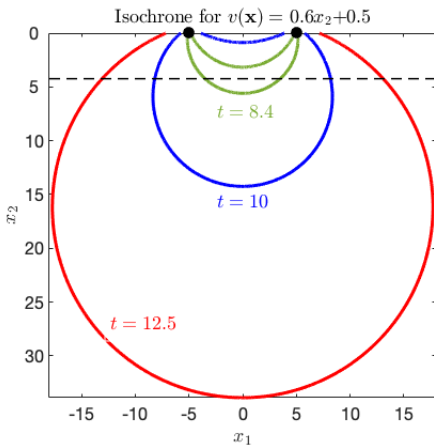
- ▶ The “isochrone” $\mathcal{J}(s, t_{\min})$ is the trivial isochrone—the geodesic between the source and receiver.
- ▶ For points with $x_2 < x_{\min}$ the operator F is not well-behaved. . .



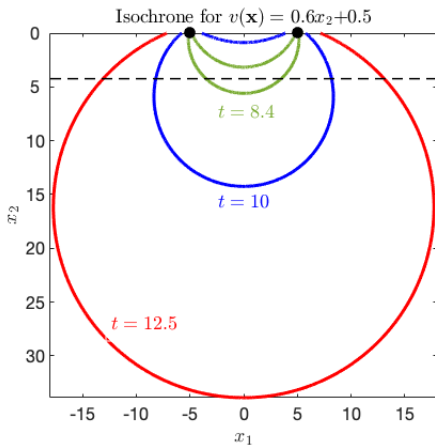
Isochrones $\mathcal{L}_{0,t}$ for $t \in \{8.4, 10, 12.5\}$ for wave speed $v(\mathbf{x}) = 0.5 + 0.6x_2$.



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Warning: Points with $x_2 < x_{\min}$ can intersect the “top” of small isochrones, causing added artifacts in the normal operator (mirror points).

F_α as a FIO with affine wave speed and $\alpha \geq 0$

$$\varphi_\alpha(\mathbf{s}, \mathbf{x}) = \tau((\mathbf{s} - \alpha, \mathbf{0}), \mathbf{x}) + \tau(\mathbf{x}, (\mathbf{s} + \alpha, \mathbf{0})),$$

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- ▶ *Canonical relation:*
 $\mathcal{C} = \{(\mathbf{s}, t, -\omega \partial_{\mathbf{s}} \varphi_\alpha, \omega; \mathbf{x}, \omega \partial_{\mathbf{x}} \varphi_\alpha) \mid t = \varphi_\alpha(\mathbf{s}, \mathbf{x}), \omega \neq 0\}$

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Why we care: When F_α satisfies the Bolker condition, $KF_\alpha^\dagger\psi F_\alpha$ is a Ψ DO, so it has many properties of differential operators, e.g., $\text{WF}(DF_\alpha^\dagger\psi F_\alpha f) \subset \text{WF}(f)$.

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Conjecture: For *some?* positive offsets, $\alpha > 0$, and affine wave speed, we conjecture that F_α satisfies the Bolker assumption.

The answer from [KQR 2023]:

Close enough phase functions give similar operators: Let F_j , $j = 0, 1$ be FIO from $\mathcal{D}'(X)$ to $\mathcal{E}'(S')$ with nondegenerate phase functions

$$\Phi_j = \omega(t - \tau_j(\mathbf{x}_s(s), \mathbf{x}) - \tau_j(\mathbf{x}, \mathbf{x}_r(s))) = \omega(t - \varphi_j(s, \mathbf{x})).$$

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Proof: If α is close to zero, then phase function ϕ_α is close to ϕ_0 .

Using the symbol to get a reconstruction operator

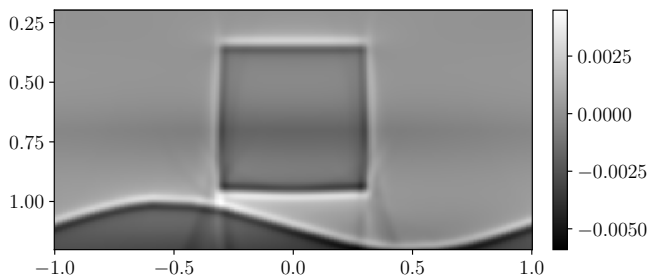
For affine velocity, we develop a reconstruction operator (like [GKQR 2018])

$$\Lambda(f) = \Delta F_0^\dagger \psi F_0 f$$

where ψ is a cutoff so F_0^\dagger and F_0 can be composed and F_0^\dagger is a backprojection with weight of F_0^\dagger adjusted to provide a more uniform symbol in depth and so a reconstruction with intensity less dependent on depth.

Here is a reconstruction where the data are generated using the wave equation, not Radon data. Unperturbed sound speed is

$$v(x_2) = 1.0 + 0.5x_2.$$



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Thank you for your attention!